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c/o Department of Geology,
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The Baram Line in Sarawak: Comments on its anatomy, history and implications for potential non-conventional gas deposits

FRANZ-LUITPOLD KESSLER

Department of Mathematics and Sciences
Curtin University of Technology, Miri, Sarawak

Abstract — This mainly fieldwork-based study reveals the Baram Line's complex history. The lineament acted as a tectonic discontinuity that joins the relatively stable Luconia Block to the relatively mobile Baram Delta Block. It acted as a boundary for both extension and compression occurring in the Baram Delta and further East in Brunei and Sabah. The more rigid Luconia was less affected by both extension and compression, whilst the neighboring Baram Delta-Sabah Foredeep crust was first stretched and later recompressed. The Baram Line may also be of interest from a gas prospector's standpoint. Reverse faulting and overthrusting is likely to have occurred in several portions of the lineament, hence creating traps. Presence of active gas charge is documented, and a potential for non-conventional gas present in slightly more porous and permeable intercalations of siltstone within the shale-dominated Setap Fm is recognized and discussed.

Keywords: Baram Line, Sarawak, tectonism, compression, strike-slip, Sundaland

INTRODUCTION

The area of studies in Sarawak encompasses coastal areas South of Miri (red rectangle in Figure 1). The Baram Line (Figure 2, Hutchison 1989) constitutes one of the most prominent lineament in Sundaland and can be traced some 140 km on the NW Borneo shelf, and some further 125 km in the molasses foreland between the coastline and the island's rugged interior. It forms an S-shaped boundary between the Luconia Block and the NE adjacent Baram Delta Block. The Southern onshore extension of Luconia is sometimes also called Balingian, and/ or Tinjar province. In this paper Luconia is referred

as to the entity West and South of the Baram Line. A comprehensive description of the Baram Delta is provided by Tan *et al.* (2000), and also Tate (2001). The Baram Line is seen as a boundary of geothermal gradient, and of crustal thickness (Hutchison 1989, 2007), as well as of hydrocarbon properties (Luconia dominated by gas, the Baram Delta by oil). Questions remain what kind of tectonic boundary the Baram Line really is; how local changes of tectonics can be explained; and how segments of this lineament may have functioned in tectonically different ways during Neogene periods.

FOLLOWING THE BARAM LINE FROM THE SEA TO THE ISLAND'S INTERIOR

In the South China Sea, the eastern edge of Central Luconia is formed by the Baram Line. On seismic as well as on gravity (Figure 3) there is little indication of deformation in the offshore part of Central Luconia. In the onshore portion, the amount of deformation is always stronger on the Baram Delta side, a mosaic of blocks which show different patterns of compressive deformation (Figure 4). As displayed in the block diagramme of Figure 5, the Luconia Block forms always the footwall, whilst the Baram Delta blocks constitute the hangingwall, with elements of inversion, folding, and thrusting. Given that portions of the onshore Baram Delta are topographically elevated (Bukit Lambir, Miri Hill, Marudi area), we are hence seeing a typical example of relief inversion.



Figure 1: Overview map of Sundaland with area of studies. From Google Earth 2009.

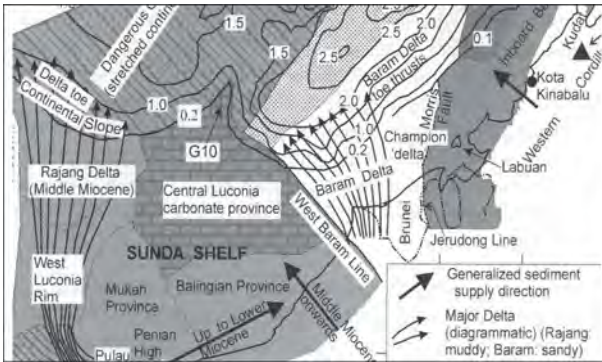


Figure 2: Map illustrating tectonic and sedimentary drivers in NW Sarawak, Brunei and Sabah (from C.S.Hutchison, 1989). The Baram Line (“West Baram Line”) is shown here to separate the stable Central Luconia carbonate shelf from the siliciclastic provinces of the Champion Delta and Baram Delta. The Baram Line originated in Oligocene - Early Miocene, when the Baram-Sabah Foredeep opened, creating a large monoclinial half-graben system, subsequently filled with Neogene clastics. During that period the Baram Line constituted the stable edge of a carbonate platform.

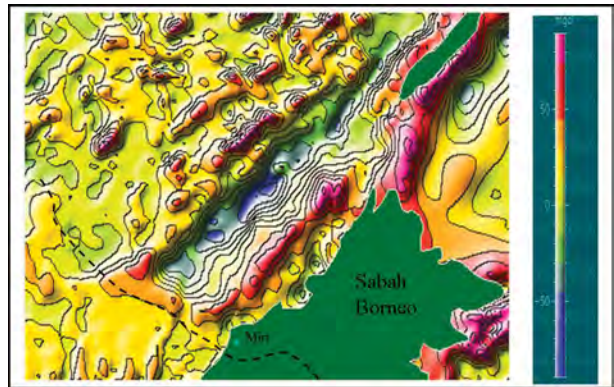


Figure 3: Map (extraction of a web-based picture, DEOS Indian Ocean Gravity project, TU Delft), illustrating Bouguer gravity west of the Borneo coast. Note the marked gravity contrast when comparing Luconia (South-West of dashed line) with the area adjacent NE, where gravity anomalies of up to -65 mgal are observed, possibly identifying an area of crustal stretching, and little compacted sediments. The alignment of gravity highs parallel to the Borneo coast further inboard can possibly be correlated to overthrust belts and clay ridges.

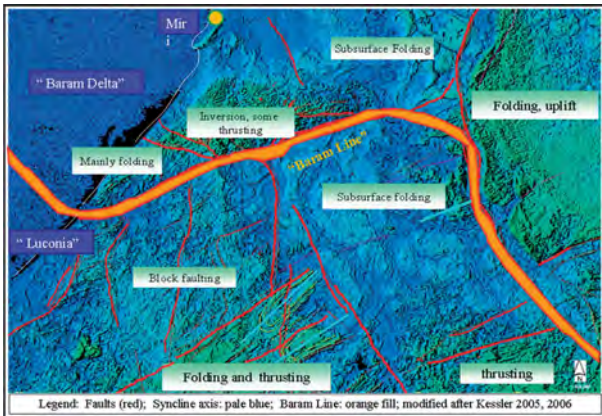


Figure 4: Northwestern Borneo can be described as a mosaic of blocks that display different style of deformation (Kessler 2005, 2006).

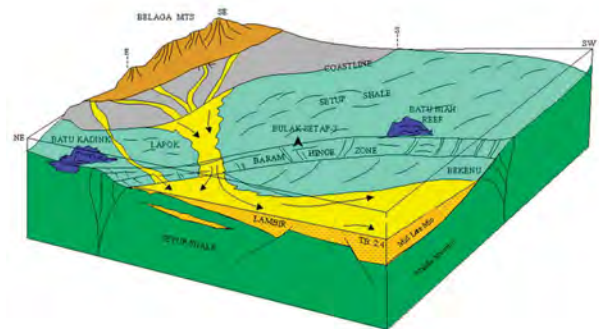


Figure 5: From a regional perspective, Central Luconia blocks are always on the footwall side, the Baram- and Champion deltas on the hanging wall facing the lineament. Given the folding of areas such as the Belait syncline, Bukit Lambir, Miri etc, NW Sarawak serves as an example of relief-inversion.

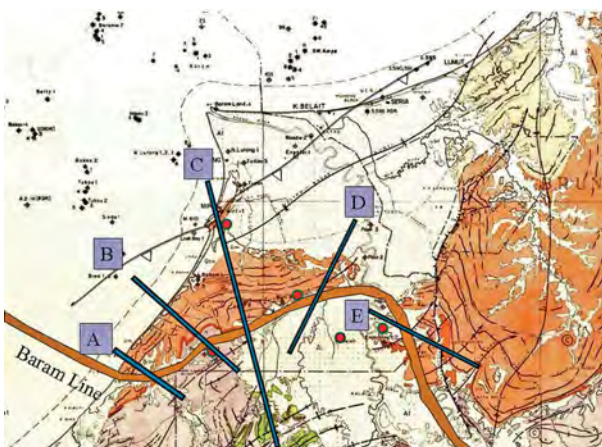


Figure 6: Geologic map of a section of NW Borneo modified after Liechti et al (1960). The map shows the location of five geologic profiles (A, B, C, D, E), and several seep locations (full circles).

Figure 6 shows a geological index map with five geological cross-sections, wells and hydrocarbon seep locations. Delineating the course of the lineament, we start at the Pantai Bungai (Bungai Beach), some fifty kilometers Southwest of Miri. Here, the strike of the line changes from NW-SE offshore to E-W direction. In Pantai Bungai, as shown in Figure 7, the Lambir anticline is seen plunging in western direction, with fossiliferous Sibuti beds outcropping at the beaches surf zone – plunging anticlines with an axial strike direction shifting from SW to NW being a common feature for many Baram Delta oil fields in the vicinity of the lineament (Hutchison 2007, p. 114).

Figure 8 shows the geological cross section “A” that runs from the Pantai Bungai toward Bukit Peninjau. As we follow the line further to the east, the hangingwall (Baram Delta) is always expressed as a high shoulder (Figure 8, Table 1), which is suggestive for inversion

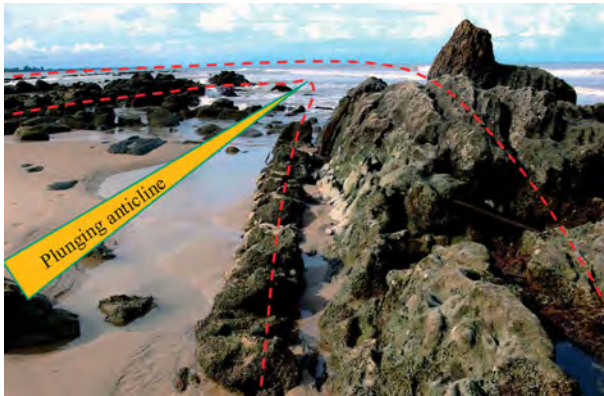


Figure 7: Bungai Beach, at seasonal low tide. Fossiliferous beds (Sibuti Fm.) rich in crab burrows and foraminifera are seen here to contour the plunging axis of the southwestern-most extension of the Lambir system.



Figure 9: Section “A” leading from the Bungai Beach, where Sibuti beds are found to dip almost vertically, towards Bukit Peninjau, where Setap Shale is only moderately folded.

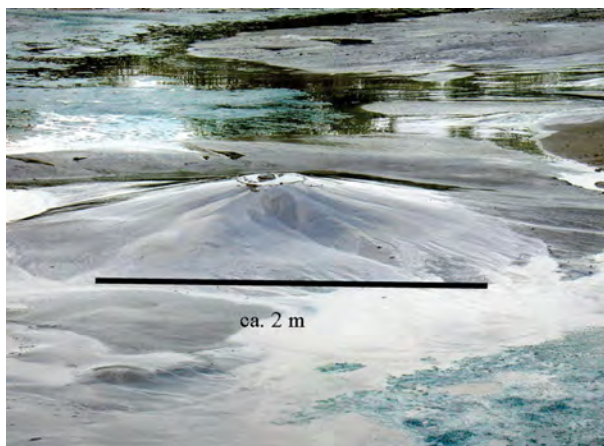


Figure 11: Section “B” leading from the Pehliau Beach, where Sibuti beds are found to dip almost vertically, towards the Bulan Setap 3 area, where gas emanates from the Ngebol mud volcano and also from a leak of a corroded wellhead. Both natural and accidental seepages document the presence of gas – entrapped stratigraphically or tectonically in Setap shale and siltstone members.

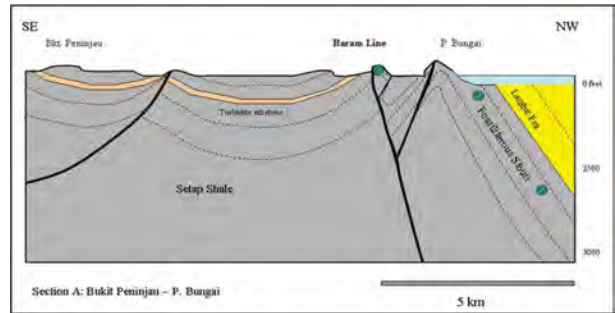


Figure 8: View of the Baram Line at the road leading to the Pantai Bungai market, near to the start of the “A section” in figure 8. The Hanging wall is formed by fossiliferous Setap Shale, the Sibuti member. Possibly the original rock wall has been shaved off during road construction work.

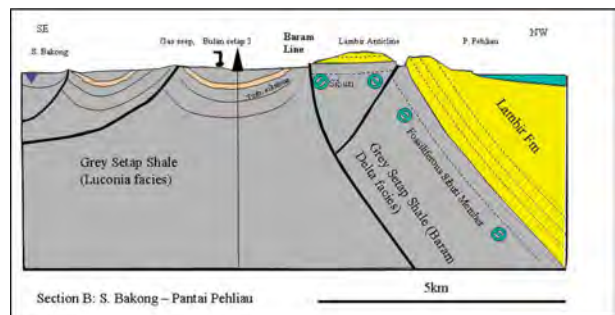


Figure 10: The Ngebol mud volcano, located only a few kilometers South of the Baram Line emanates mud and gasses at shifting rates. Flammable gas is also seeping-off from the corroded wellhead of Bulan Setap 3 (Sawarak Oilfields Ltd, abandoned in 1935).

occurring along the fault, with the strong possibility of over-steepening and faultplane-rotation from vertical to reverse-fault or even to overthrust angle. The hangingwall section is characterized by minor faulting with fold axis perpendicular or oblique to the Baram Line. This is also a typical feature on the Kawang thrust sheets of the Miri Hill, where several folds oblique to the thrusts are observed.

In cross-section “B” (Figure 9) we observe a mildly deformed Luconia Foreland, with grey Setap Shale at surface. This Setap section contains a sequence of siltstone-fine sandstone turbidites. These are not seen on the Baram Delta side, which is dominated by the fossiliferous Sibuti Fm. The section “B” also crosses the Bulak Setap 3 location, and the near-by mud volcano of Ngebol (Figure10).

Several kilometers further East, section “C” crosses the highest areas of both Bukit Lambir and also, further NE, the Siwa-Miri-Seria overthrust (Figure 11). From here onwards, the Baram Line now roughly follows the course of the Sungai Bakong, occasionally incising into the hangingwall by up to 20 meters. Section “D” (Figure 12) characterizes an area of less deformation compared to the previous “C”, as the Lambir/Belait beds are seen dipping monotonously in NE direction.

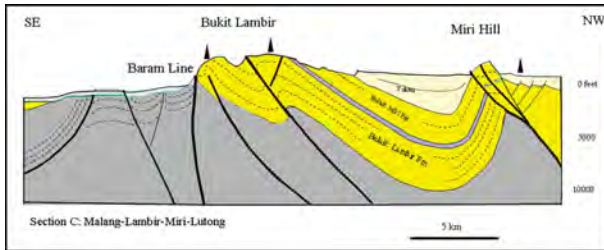


Figure 12: Section “C” leads from Miri to Lambir, and the Beluru area. Both Lambir and Miri are shaped by folding, and overthrusting. The overthrust component in the Baram Line is difficult to judge in absence of reliable subsurface data. Entrapment of non-conventional “shale” gas in the hanging wall below Lambir could be considered in presence of proven gas charge.

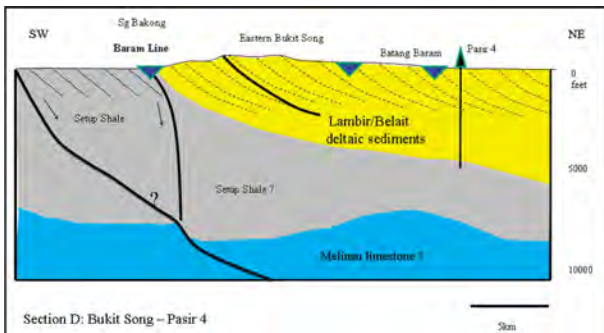


Figure 13: Section “D” features Eastern Lambir (Bukit Song). The style of deformation appears less strong if compared with “C”.

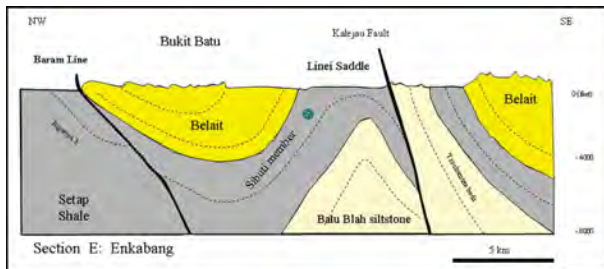


Figure 14: Section “E” runs from the Belait escarpment, the Kalejau fault system into the foothill area of Bukit Enkabang. Similar to Section “C”, the amount of overthrusting by the Baram Line might be significant and offer similar entrapment possibilities as quoted for section “C”. Modified after Kessler 2005, 2006 and partly based on SOL data generated by Knecht (1928), Whooley (1928), Morgan (1931) and Avang Damit (1940).

Near to the confluence point of the Sungai Bakong with the Batang Baram, the Baram Line is taking a turn from hading East to South-East. From there it also runs parallel with another fault zone, the Kalejau fault, a few kilometers afar. The area around marudi is characterized by intense block-faulting and block-rotation, with block sizes in the order of several square kilometers. As shown in Section “E” of Figure 13, there could be an important element of overthrusting along the faults, and several oil/gas seeps have been reported by SOL geologists in unpublished reports. From the Marudi area further

South-East, the Baram Line fades on with the same strike direction, and can be traced until Long Lama. There it merges near Batu Gading with overthrusts at the suture of the foreland with the Rajang mountain belt. Further fieldwork seems required to delineate the path of the Baram Line.

TECTONIC HISTORY OF THE BARAM LINE

Given that the Baram Line is strongly associated with the formation of the Baram Delta, it could be argued that the lineament originated as early as the late Oligocene. Recent work by Simons *et al.* (2007) also suggests that at least some segments of the Baram Line are still actively moving at present (Figure 15). We are hence dealing with a lineament that has relatively old origins (Oligocene), but also a very recent overprint. Accordingly, one might expect complexity when describing the movements along the Baram Line as vectors of extension and compression may have changed during the Neogene.

In simplified terms, one could describe the history of the lineament in three phases: Extension plus strike/slip; compression; compression plus strike/slip but in an opposite direction.

Figure 16 describes the earliest part of the history: East of central Luconia, the crust is being stretched and thinned, which lead to the formation of the Baram-Sabah Foredeep, and an associated rapid subsidence. The rise of the (Rajang Group) Belaga Mountains, occurring at the onset of the Middle Miocene, provided massive amounts of clastics that filled up the new depocenters.

Compression occurring at the late Miocene/Pliocene boundary caused inversion, folds and thrust-sheets, and the stretched crust underlying the Oligo-Mio graben system may have buckled and shrunk by some estimated 1 to 5 % in width (Figure 17).

With ongoing compression, but crustal consolidation being realized, the Baram Line was remobilized as a strike-slip system (Figure 18). Studies encompassed by Simons *et al.* (2007) indicate active movements against or along the Baram Line in the order of 3 mm/y, but the existing data are too sparse to judge, if the stress is relieved by the Morris Fault/Jerudong Line, the Baram Line, or both systems combined.

DISCUSSION

Remains the question where the 3mm/year of crustal shortage are accommodated. The most simple answer is that most of it could be released in deep reverse faulting/overthrusting within the Baram Delta- Sabah Foredeep, and the Bouguer map shown in Figure 2 suggests this might have happened, though requiring further confirmation by deep seismic data and perhaps also geodetic measurements. A more farfetched possibility is to bring the Baram Line in context with the Sulawesi subduction zone and possibly the Red River Fault system of China and Indochina.

Regarding the latter, conclusive evidence for genetically linking those two intra-continent fault systems (Baram Line, areas of Sundaland to Sulawesi; Red River Fault, east of the Himalayas) has not been provided.

Economic aspects: Potential for non-conventional gas deposits

Previous exploration campaigns carried out in the onshore Baram Delta mostly targeted deltaic topset reservoirs of the Belait Group. In most areas reservoir were found too shallow and too poorly sealed to yield substantial petroleum deposits – exceptions being the Miri Field, Rasau/Asampaya at the Brunei border and the very large Seria in Brunei. The quest for deeper clastic reservoirs of good quality has so-far been disappointing.

Regional charge

From regional drilling results and seep data it is known that two charge systems seem to co-exist side-by-side: A predominantly oil charge system for the Baram Delta area, and predominantly gas for the Luconia area. This view, however, excludes the potential differential effect of sealing capacity. Deeper (and potentially more gassy) levels of the Baram Delta remain poorly explored to date. Post-WW II exploration in the Miri Field has pointed toward the existence of deeper reservoirs and a gas potential.

Oil and gas window

In respect to the depth of oil and gas generation, vitrinite reflectivity measurements (VRM) were carried out. Samples from Lambir/Belait sediments above the Grey Setap Shale on the Luconia side of the lineament, as well as in the middle-to-upper section of Lambir sandstone gave values of VRM 0.42 and 0.43 (Kessler 2008, in print). These values are suggestive for a removal of sediments in the molasses foreland area South of Miri in the order of some 2000 feet. This would put the top of economic oil generation at ca. 7000 feet below surface, and top gas generation at some 11000 feet. These numbers are based on an estimated temperature gradient of 3 deg C/ 100 m.

Evidence of charge

In respect of hydrocarbon charge, there is plenty of evidence that charge is active on both sides of the Baram Line, and gas may be contained in marginal siltstone reservoirs such as turbiditic siltstones in the Bukit Peninjau area, the Batu Blah siltstone South of Marudi. Expected poroperm values at target depth might be in the order of 8 % porosity and less than 10 md permeability (in comparison with regional data). To date it is not known if the Setap Shale Fm *sensu stricto* can be a host rock for gas.

Trap characteristics

The size of entrapment may be correlative to the amount of inverse faulting/overthrusting along sectors

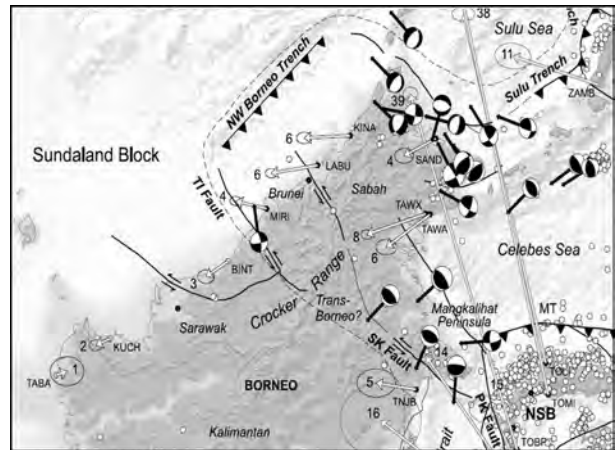


Figure 15: According to GPS measurements, (Simons et. al. (2007)), conjugated pair of Jerudong and Baram Line form a boundary between relatively stable SW Sundaland crust, and the more ductile and possibly thinner Eastern Borneo. Measurements are suggested of E-W compression, and compression/absorption of crust in the order of 3 mm/year between SW and NE Borneo. These data may explain the presence of Late Miocene and Pliocene overthrusting as well as movements along lineaments in strike-slip mode.

of the Baram Line, with sealing provided by clay and clay gouging along the fault plane. Typically tectonics as described above lead to elongated lenticular-shaped traps. Most likely the lateral amount of overthrusting may be in the order of hundreds of meters, or perhaps several kilometers but this cannot be certified without good quality seismic data.

Play summary

A play section is shown in Figure 19. The question remains how much recent tectonism is at work, having possibly affected gas traps in negative ways. This said, non-conventional gas is less sensitive to breach-and-spill compared to gas located in high porosity/permeability reservoirs. Using modern high-resolution and high-fold 2D seismic the overthrust-related traps can easily be imaged.

CONCLUSION

The study, mainly based on fieldwork, reveals the Baram Line's complex history. The lineament acted as a tectonic discontinuity that joins the relatively stable Luconia Block to the relatively mobile Baram Delta Block, acting as boundary for both extension and compression occurring in the Baram Delta. A potential for non-conventional gas in the Setap Shale and/or slightly more porous and permeable intercalations of siltstone is recognized along segments of the lineament.

ACKNOWLEDGEMENTS

The author would like to thank Prof. Dr. Charles S. Hutchison, Andrew B. Cullen (Shell), M. Wiemer (Shell), Prof. Dr. Chai Peng Lee and Dr. Eswaran Padmanabhan for discussions.

Table 1: Calibration points for the course of the lineament with GPS coordinates.

Area	Northing	Easting	Altitude (m)	Comment
Bungai Beach	04°03.733'	113°47.041'	2.5	50 m east of market, high shoulder (hanging wall)
Bridge, Sg. Jangalas	04°03.890'	113°49.918'	5	Junction to Kampong Menjelin
Kg. Menjelin road	04°04.752'	113°50.466'	6	Several meters of high shoulder (hanging wall)
Kg. Kejapil Beraya road	04°08.582'	113°54.290'	18	Relief contrast (hills on Lambir side)
Old Miri-Bintulu trunk road	04°08.772'	114°00.325'	52	Strong relief contrast (rise towards Bt. Lambir)

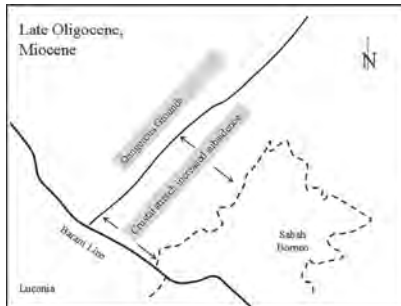


Figure 16: With the opening of the Baram Delta – Sabah Foredeep (compare Tan et al 2000), the Baram Line may have acted as a transform fault for a while. Little deformation is quoted from Luconia areas opposite to the opening foredeep.

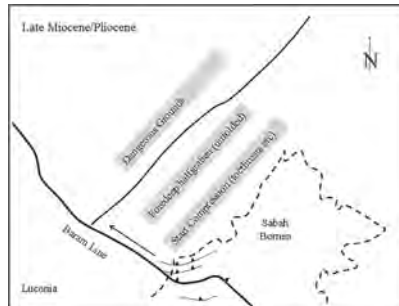


Figure 17: Compression affected the area during the late Miocene to Pliocene period (Kessler 2005, 2006), leading to inversion, folding and some thrusting features.

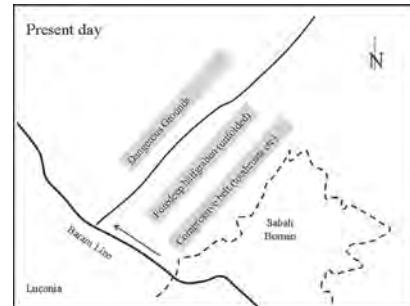


Figure 18: Present day measurements are suggestive for strike-slip movements along the Baram Line; compressive stress may be accommodated by toethrusts and mobile clay ridges from Sarawak to Labuan island

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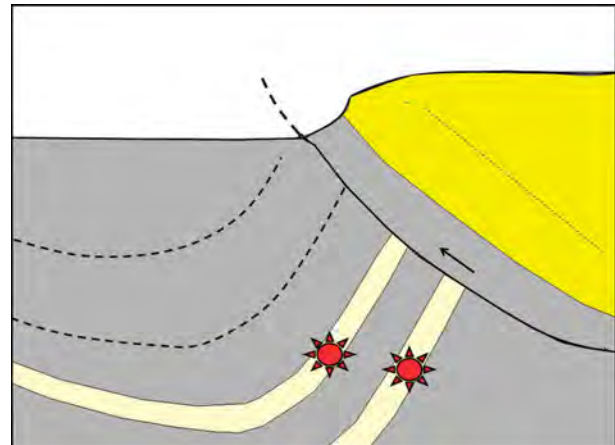


Figure 19: Play section for reverse fault / overthrust segments of the Baram Line. Depending on the fault angle, there is an entrapment potential in the footwall section; seep data are suggestive that hydrocarbon charge is still being generated, whereas reverse faulting/thrusting may be confined to the Pliocene.

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Penemuan bivalvia *Daonella pahangensis* Kobayashi berusia Trias Tengah dari Aring, Kelantan

AHMAD ROSLI OTHMAN^{1,2,*} & MOHD SHAFEEA LEMAN²

¹Jabatan Mineral dan Geosains Malaysia,
Wisma Persekutuan, 15200 Kota Bharu, Kelantan

*Email address: ahmadrosli@jmg.gov.my

²Institut Alam Sekitar dan Pembangunan (LESTARI)
Universiti Kebangsaan Malaysia, 43600 Bangi, Selangor

Abstrak— Satu spesimen Bivalvia *Daonella pahangensis* Kobayashi telah ditemui pada satu singkapan batuan di Lokaliti QZ467 dalam Formasi Telong di kawasan Aring, Gua Musang. Ini adalah penemuan pertama spesies tersebut di Kelantan selain yang telah ditemui dan diperihalkan oleh Kobayashi (1964) di Temerloh, Pahang. Spesies ini tergolong dalam Kumpulan *Daonella sturia-Daonella lommeli* yang bersifat endemik dan terhad dalam rantauan Tethys. Spesies bivalvia ini berasosiasi rapat dengan *Posidonia* sp. dan wujud bersama ammonoid *Frankites regoledanus* (Mojsisovics) yang merupakan penanda subzonal bagi Subzon *Regoledanus* yang berusia Ladinian akhir, Trias Tengah.

The discovery of Middle Triassic Bivalve *Daonella pahangensis* Kobayashi from Aring, Kelantan

Abstract— A single specimen of bivalve *Daonella pahangensis* (Kobayashi) has been discovered from an outcrop at Locality QZ467 within the Telong Formation, in the vicinity of Aring, Gua Musang. This is the first discovery of such species in Kelantan besides those described by Kobayashi (1964) in Temerloh, Pahang. This species is commonly included in *Daonella sturia-Daonella lommeli* Group, an endemic group and restricted to the Tethyan region. This bivalve species is closely associated with *Posidonia* sp. and it often co-exists with late Ladinian ammonoid *Frankites regoledanus* (Mojsisovics) which is a subzonal marker for *Regoledanus* Subzone of late Ladinian, Middle Triassic.

Keywords: Bivalve, *Daonella pahangensis*, Aring, late Ladinian, *Regoledanus* Subzone

PENGENALAN

Fosil *Daonella* adalah biasa didapati dalam batuan sedimen Trias Tengah di Malaysia. Berdasarkan kewujudannya, Kobayashi (1964) membahagikan strata batuan sedimen kepada dua zon iaitu zon di sebelah barat meliputi kawasan Kedah hingga ke Perak dan zon timur yang meliputi kawasan selatan Kelantan, Pahang dan Johor. *Daonella* dalam zon di sebelah barat berasosiasi rapat dengan *Halobia* dan *Posidonia*. Manakala *Daonella* dalam zon di sebelah timur pula dicirikan dengan kehadiran batu pasir *Myophoria* dan himpunan ammonoid yang berusia Anisian.

Laporan ini mendokumentasikan penemuan fosil Bivalvia *Daonella pahangensis* Kobayashi dalam strata batuan sedimen berusia Trias di Aring, Gua Musang Kelantan yang terletak dalam zon di sebelah timur.

KAJIAN TERDAHULU

Fauna Bivalvia Trias *Daonella* telah banyak dilaporkan penemuannya di kawasan Perak, Kelantan, Pahang dan Kedah oleh beberapa pengkaji terdahulu. Jones *et al.* (1966), Jones telah melaporkan bahawa Kobayashi merekodkan penemuan *Daonella indica* Bittner di Kuala Nerang, Kedah. Manakala Kobayashi (1964) dan Jaafar Ahmad (1976) telah melaporkan penemuan beberapa spesies *Daonella* yang telah ditemui oleh Procter di beberapa lokaliti di Temerloh, Pahang. Spesies-spesies tersebut termasuklah *Daonella indica* Bittner, *Daonella sakawana* Mojsisovics, *Daonella pahangensis* Kobayashi, *Daonella pichleri* Mojsisovics dan *Daonella lommeli* Wissman. Metcalfe *et al.* (1982), menemui *Daonella lommeli* Wissman dan *Daonella sakawana* Mojsisovics di kawasan Mentakab, Pahang. Di Kelantan, Kobayashi *et*

al., (1966) merekodkan penemuan *Daonella* sp. di Sungai Lebak manakala Yin (1963) pula menemui *Daonella* cf. *indica* Bittner di Sg. Chiku, Gua Musang.

Berdasarkan penemuan *Daonella* tersebut, didapati *Daonella* yang ditemui di kawasan Aring, Gua Musang, Kelantan memperlihatkan morfologi yang menyamai spesimen *Daonella pahangensis* Kobayashi.

GEOLOGI

Fosil ini ditemui dalam lapisan batu lumpur kelabu yang sebahagiannya bertuf. Batu lumpur ini berlapis nipis dan telah termiring, dengan jurus sekitar 340°-350° dan kemiringan 25°-30° ke timurlaut. Kawasan penemuan fosil terletak pada koordinat RSO: QZ 467902 WMR: 536601 yang dirujuk sebagai Lokaliti QZ467. Kawasan kajian berada dalam Jajahan Gua Musang dan terletak di bahagian tenggara Negeri Kelantan (Rajah 1). Berdasarkan pemetaan geologi yang dilakukan oleh Aw (1990) di kawasan Aring, kawasan ini terletak dalam Formasi Telong yang setara dengan Formasi Semantan.

USIA DAN TABURAN

Fosil Bivalvia *Daonella pahangensis* Kobayashi telah ditemui bersama fosil-fosil ammonoid yang berusia Ladinian akhir iaitu *Frankites regoledanus* (Mojsisovics), *Protrachyceras costulatum* Mansuy, *Daxatina* sp., *Protrachyceras* cf. *pseudoa-archelaus* Mojsisovics, *Anolcites* sp., *Joannites* sp., *Clionitites* sp., *Zestoceras* sp. dan *Megaphyllites jarbas* Mojsisovics oleh Ahmad Rosli & Mohd Shafeea (2008). Ammonoid *Frankites regoledanus* (Mojsisovics) adalah merupakan ammonoid yang biasa terdapat dalam laut Paleo-Tethys serta merupakan penanda subzon bagi Subzon *Regoledanus* dalam Zon *Protrachyceras* (Mietto & Maftrin, 1995). Spesies yang bersifat endemik ini juga bersekutuan dengan Bivalvia

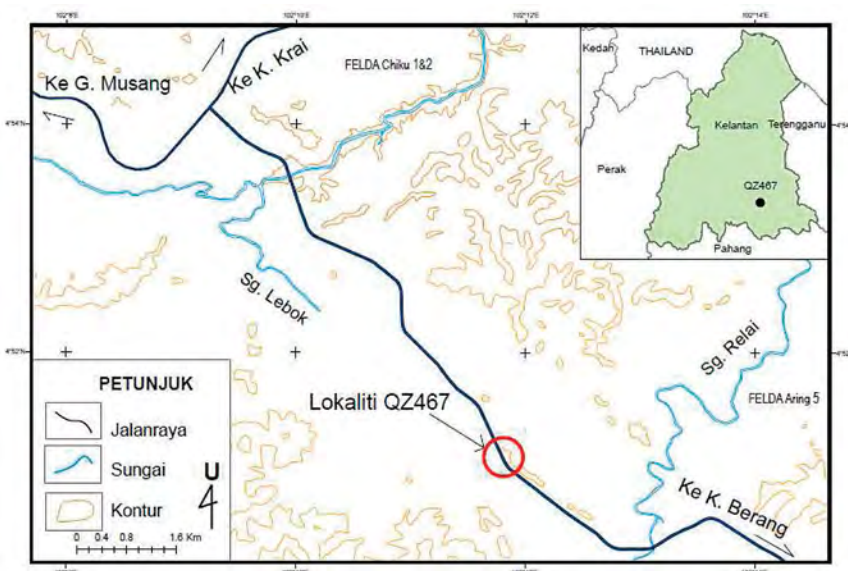
Posidonia sp. Daripada kumpulan-kumpulan spesies *Daonella* di Malaysia yang diutarakan oleh Kobayashi (1964), spesies ini tergolong dalam Kumpulan *Daonella sturia*-*Daonella lommeli*.

Halobiidae (Bivalvia) iaitu *Daonella* dan *Halobia* adalah bertaburan meluas di lautan Tethys, Panthalassa dan Artik sepanjang zaman Trias Awal ke Akhir iaitu Anisian hingga ke Norian. *Daonella* merupakan pincirian utama terhadap ekologi laut dalam yang tahan terhadap tekanan tinggi dan persekitaran yang kurang oksigen (Kobayashi *et al.*, 1966). Turut sama dalam ekologi tersebut termasuklah spesies *Posidonia* dan *Halobia* dan selalunya bersekutuan dengan ammonoid nekton. *Daonella* lebih berkembang semasa Ladinian manakala *Halobia* pula berkembang semasa Karnian (McRobert, 2000; Kobayashi & Tokuyama, 1959).

Menurut McRobert (2000) Bivalvia ini adalah kosmopolitan dan sering ditemui dalam fasies syal hitam dalam persekitaran anoksik atau disoksik. Menurutnyanya lagi dari segi poligenetik *Halobia* mungkin bukanlah taksa semulajadi tetapi adalah polipiletik (polyphyletic) daripada keturunan *Daonella* atau posidoniid lain. Menurutnyanya juga, Bivalvia *Halobia daonellaformis* McRobert yang ditemui di British Columbia, Kanada merupakan spesies peralihan antara *Halobia* dan *Daonella*. Menariknya, tidak seperti *Posidonia*, spesies-spesies *Daonella* jarang didapati wujud bersama dalam satu lapisan. Ini menimbulkan persoalan adakah keadaan ini bergantung kepada perbezaan kronologi atau keadaan ekologi (Kobayashi & Tokuyama, 1959).

MORFOLOGI

Daonella dapat dibezakan dengan *Halobia* melalui perbezaan cuping aurikel anterior. Bagi *Daonella* cuping aurikel mempunyai rib yang berpusat pada umbo tetapi *Halobia* pula membentuk segi tiga pada cuping aurikel



Rajah 1: Rajah menunjukkan lokasi penemuan fosil Bivalvia *Daonella pahangensis* Kobayashi dan kedudukan Lokaliti QZ467.

tanpa rib atau tidak mempunyai cuping aurikel. Morfologi kedua-duanya digambarkan seperti dalam Rajah 2.

Paleontologi Sistematik

Superfamili PECTINOIDEA Rafinesque

Famili POSIDONIIDAE Frech

Genus *Daonella* Mojsisovics

Daonella pahangensis Kobayashi, 1964

Plat 1, Gambar 1 dan 2

1964 *Daonella pahangensis* Kobayashi- Kobayashi: ms 62, Plat V, gambar 11.

Pemerihalan: Cengkerang bersaiz besar dan pipih, inequilateral. Ukuran panjang melebihi tinggi cengkerang. Garis engsel memanjang lurus, lebih pendek daripada panjang cengkerang. Umbo terletak kira-kira setengah daripada garis engsel. Rib berjejeri lurus dengan sedikit terserong berhampiran umbo. Rib primer lebih kurang 25 garis menjadi cabang dua berhampiran umbo dan wujud cabang tiga pada jarak satu perdua dari umbo tetapi tidak melebihi bilangan garis cabang dua. Garis pertumbuhan terpusat agak jelas pada jarak satu pertiga melingkari umbo.

Bahan: Satu spesimen cengkerang kanan diperiksa yang mewakili sampel QZ467/B1. Bahagian postero-ventral cangkerang tidak dapat dicerap. Oleh itu ukuran dimensinya adalah berdasarkan anggaran sahaja.

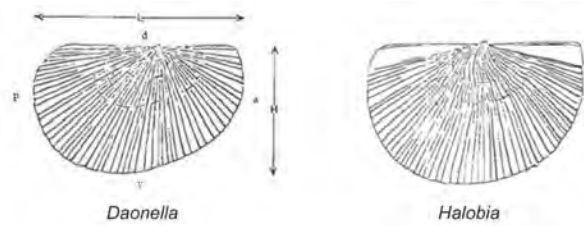
Dimensi:	Tinggi(H)	Panjang (L)
QZ467/B1 (cangkerang kanan)	± 38 mm	± 60 mm
Holotip (cangkerang kiri)	45 mm	70 mm

Pernyataan : *Daonella pahangensis* Kobayashi merupakan spesies bersaiz paling besar daripada spesies-spesies *Daonella* yang lain. Dari segi dimensi dan perhiasan cangkerang, spesies ini menyamai spesimen holotip yang ditemui di kawasan Temerloh oleh Kobayashi (1964). Spesies ini juga menyerupai *Daonella sakawana* Mojsisovics dari segi garis bentuknya tetapi lebih quadrat serta lebih kecil dimensinya berbanding *Daonella pahangensis* Kobayashi. Selain itu garis rib cabang tiga yang wujud pada *Daonella pahangensis* Kobayashi membezakannya dengan *Daonella sakawana* Mojsisovics yang mana rib cabang dua lebih dominan dan jelas.

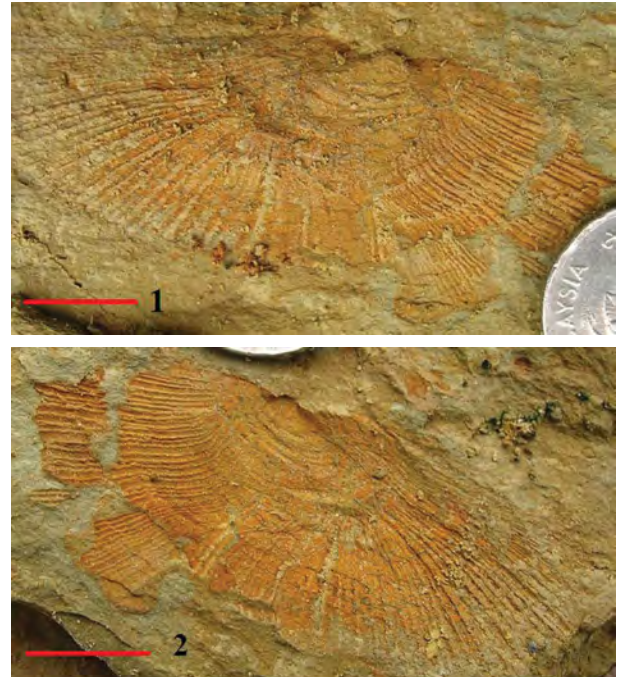
Kewujudan : Lokality QZ467, Formasi Telong, Aring, G. Musang

KESIMPULAN

Bivalvia *Daonella pahangensis* Kobayashi yang sebelum ini hanya ditemui di Temerloh, Pahang turut ditemui di Kelantan dalam turutan batuan Trias di Lokality QZ467, di kawasan Aring. Berdasarkan assosiasinya dengan fosil penanda subzon iaitu ammonoid Tethys *Frankites regoledanus* (Mojsisovics) usia pegenapan batuan enapan marin laut ini adalah pada usia Ladinian akhir, Trias Tengah.



Rajah 2: Morfologi katup kanan bagi *Daonella* dan *Halobia* digambarkan secara diagramatik. Parameter statistik H: tinggi; L: panjang; d: sisi dorsal; v: sisi ventral; a: sisi anterior dan p: sisi posterior. Dari Silberling (1963).



Plat 1: *Daonella pahangensis* Kobayashi, 1964. Skala bar mewakili 1 cm. Lokasi: QZ467, Formasi Telong, Aring, Gua Musang, Kelantan. Foto 1: Acuan dalaman, cangkerang kanan, spesimen QZ467/B1. Foto 2: Acuan luaran, cangkerang kanan, spesimen QZ467/B1

PENGHARGAAN

Penulisan artikel ini adalah sebahagian daripada hasil penyelidikan Sarjana yang sedang dijalankan oleh penulis di bawah seliaan Prof. Dr. Mohd Shafeea Leman di Universiti Kebangsaan Malaysia. Ucapan terima kasih ditujukan kepada Pengarah JMG Kelantan, Hj. Zainol Abidin Sulaiman kerana telah memberikan kesempatan dan sokongan untuk menjalankan penyelidikan ini.

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Manuscript received 29 June 2009

PERTEMUAN PERSATUAN MEETINGS OF THE SOCIETY

RETIREMENT LECTURE No. 3

Effects of the financial crisis on the global mining industry

TEOH LAY HOCK

Monday, 27 July 2009

Department of Geology, University of Malaya

Mr. Teoh Lay Hock who retired from the Minerals and Geoscience Department gave his retirement lecture on the effects of the financial crisis on the mining industry globally. The talk was attended by academic staff and students from University of Malaya and Universiti Kebangsaan Malaysia, as well as several private practitioners. This talk is the third in the series organised by Mr. Tan Boon Kong, Chairman of the Working Group on Engineering Geology, Hydrogeology and Environmental Geology.



MALAM GEOSEA RE-RUNS

Friday, 14 August 2009,
Department of Geology, University of Malaya

5.00 – 5.30 pm :	Tea
5.30 – 6.00 pm :	Sdr. Zakaria Mohamad (JMG) The Canada Hill Landslide, Miri, Sarawak
6.00 – 6.30 pm :	Sdr. Chow Weng Sum (UTP) A Case Study on Limestone Cliff in Tambun, Perak
6.30 – 7.00 pm :	Q & A

Mr. Tan Boon Kong, Chairman of the Working Group on Engineering Geology, Hydrogeology and Environmental Geology organised the re-run of two papers presented at the Eleventh Regional Congress on Geology, Mineral and Energy Resources of Southeast Asia (GEOSEA 2009). This event was organised in collaboration with the Institute of Geology Malaysia. The event was well attended by academics, engineering geologists and engineers.



CERAMAH TEKNIK TECHNICAL TALK

The pre-opening history and opening of the Black Sea

DR. RANDELL STEPHENSON

Reader in Geophysics, School of Geosciences, University of Aberdeen

Monday, 17 August 2009

Department of Geology, University of Malaya

Abstract— The Black Sea is a back-arc basin with active extension beginning in Early-middle Cretaceous times. Extensional structures within and around the basin are poorly known and are mostly overprinted (inverted) by a subsequent Tertiary compressional stage. Black Sea extension itself probably reactivated earlier extensional structures dating from at least the Permo-Triassic and probably to the Late Devonian. There is a fairly vigorous debate about the relative ages of western and eastern segments of the Black Sea basin, with the former usually considered to be somewhat older than the latter. What this could mean for the geometry of extension and mechanism(s) of back-arc rifting processes in terms of the broader plate configuration and kinematics are poorly defined. I'll summarise these issues, including the present-day tectonic regime and what can be inferred about underlying lithosphere structure, and consider these in the context of models of modern back-arc basin systems. The idea is to elucidate the processes responsible for the formation of the Black Sea and to help define a template for interpreting its under constrained observational dataset.



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3. Nuruldiyana Mohd Saharudin

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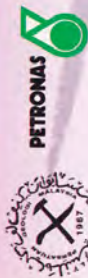


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- 15/9/2009 -- confirmation of paper selection
- 2/10/2009 -- end of exhibition form submission
- 15/10/2009 -- end of extended paper/ final paper submission
- 30/10/2009 -- announcement of exhibitors list
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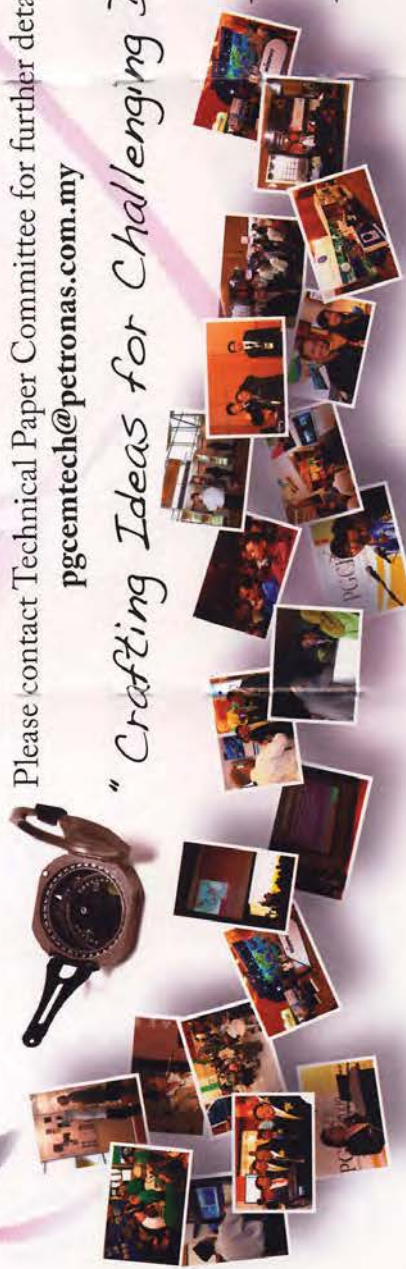
- Geophysical & Acquisition
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- Exploration Case Study
- Reservoir Characterization & Modeling
- Geosciences R & D.

****Kindly please mention the category of your paper/ poster.**

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Due date for extended abstract/ final paper submission: 15/10/2009.

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Note:

Advanced registration will be accepted until 28/2/2010.

Payment by credit card is now available through online registration, with automatic 4% service charges.

For local participants, payment by crossed cheque to be made payable to the Geological Society of Malaysia. Outstation cheques should over bank charges.

The registration fee for overseas participants can be paid by telegraphic transfer. Details are as follows:
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Please mail/ fax the remittance advice slip of the telegraphic transfer transaction.

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UPCOMING EVENTS

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March 8-12, 2010: Basic Petroleum Geology, Kuala Lumpur, Malaysia. Tel: 603 21684751; email: ap-enquiries@petroskills.com

March, 8-11, 2010: Oil & Gas Libya Expo 2010. Contact: Vijay Arjun, Dubai; Tel: 00971 4 2988144; Fax: 00971 4 2987886; email: vijay@orangeairs.com; website: www.orangeairs.com

March 16-18, 2010: Offshore Asia Conference & Exhibition 2010, Kuala Lumpur Convention Centre, Kuala Lumpur, Malaysia. Secretariat: Suthi Chatterjee, Tel: 662 247 6533; Fax: 662 247 7868; email: suthi@prmc-thailand.com; Web: www.offshoreasiaevent.com

March 29-30, 2010: Petroleum Geology Conference & Exhibition, Kuala Lumpur Convention Centre, Kuala Lumpur, Malaysia. Contact: Tel: 603 79577036; Fax: 603 79563900; email: geologi@po.jaring.my; website: www.pgcem.com

April 5-9, 2010: Basic Geophysics, Kuala Lumpur, Malaysia. Contact: Tel: 603 21684751; email: ap-enquiries@petroskills.com

April 6-8, 2010: Asiawater '10 Expo, Kuala Lumpur Convention Centre, Kuala Lumpur. Contact: Tel: 603 4045 4993

April 12-16, 2010: AVO, Inversion and Attributes: Principles and Applications, Kuala Lumpur, Malaysia. Contact: Tel: 603 21684751; email: ap-enquiries@petroskills.com

April 13-15, 2010: Offshore Australasia Conference, Burswood Convention Center, Perth, Australia. Contact: marketing@neventurecorp.com; website: www.neventurecorp.com/2010/0ac

April 14-15, 2010: Offshore Australasia Conference 2010, Perth, Australia. Website: www.neventurecorp.com/2010/0ac

April 19-23, 2010: Sequence Stratigraphy, Vienna, Austria. Tel: +43 3842 43053-33; Fax: +43 3842 430531; email: training@hoteng.com; website: www.hoteng.com

April 26-28, 2010: 1st International Applied Geological Congress, 2010, Department of Geology, Faculty of Science, Islamic Azad University – Mashhad Branch (IAUMB), Mashhad-Iran. Email: info@iagc.ir; website: http://www.iagc.ir

May 3-7, 2010: Fundamentals of Petroleum Geology, Vienna, Austria. Tel: +43 3842 43053-33; Fax: +43 3842 430531; email: training@hoteng.com; website: www.hoteng.com

May 3-7, 2010: Mapping Subsurface Structures, London, UK. Tel: 603 21684751; email: ap-enquiries@petroskills.com

May 10-14, 2010: Sandstone Reservoirs, Kuala Lumpur, Malaysia. Tel: 603 21684751; email: ap-enquiries@petroskills.com

May 10-14, 2010: Analysis of Structural Traps in Extensional Settings, London, UK. Tel: 603 21684751; email: ap-enquiries@petroskills.com

May 18-20, 2010: The 34th Indonesian Petroleum Association Convention & Exhibition 2010, Jakarta Convention Center, Indonesia. Tel: 62 21 5155959; Fax: 62 21 51402545; email: tpc@ipa.or.id; website: ipa.or.id/34th-Convention/CFP2010.htm

June 7-11, 2010: Reservoir Geology – Integrating Data for Reservoir Modelling, Vienna, Austria. Tel: +43 3842 43053-33; Fax: +43 3842 430531; email: training@hoteng.com; website: www.hoteng.com

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June 21-25, 2010: Operations Geology, Kuala Lumpur, Malaysia. Tel: 603 21684751; email: ap-enquiries@petroskills.com

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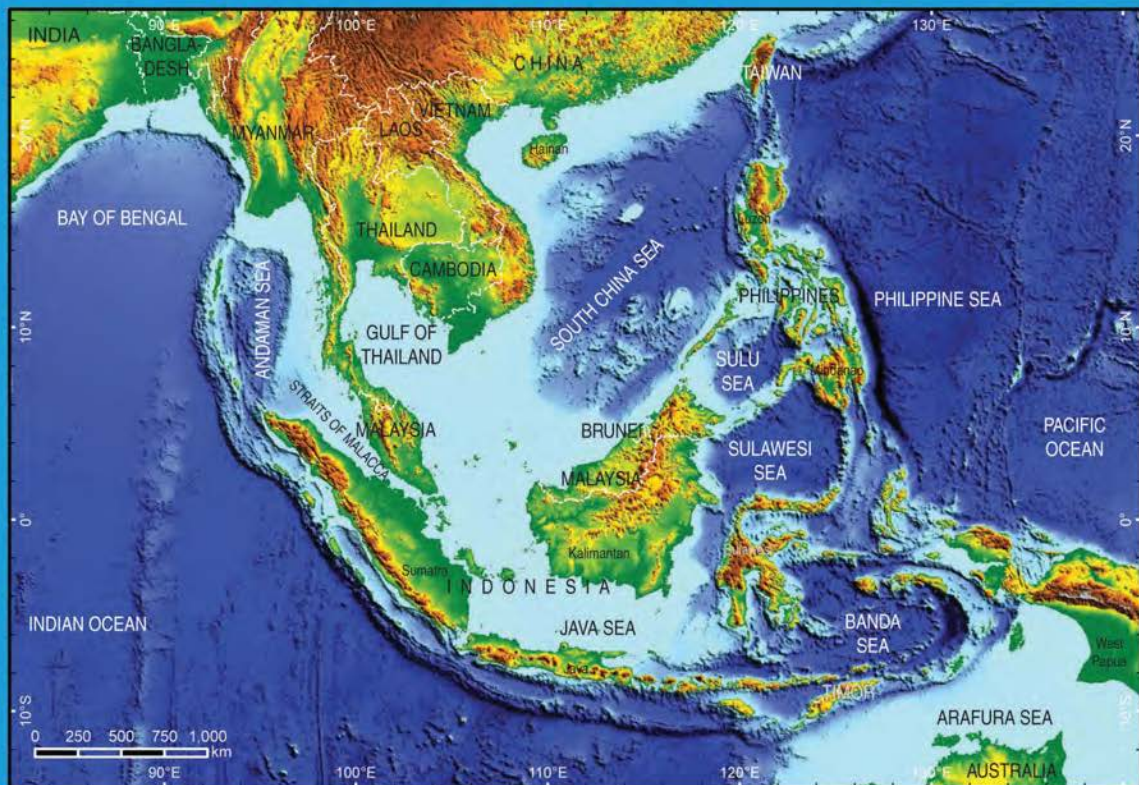
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Email: geologi@po.jaring.my

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The paper can be written in Bahasa Malaysia (Malay) or English. For English papers, use either British or American spelling but not a combination of both. The paper should be checked thoroughly for spelling and grammar. The manuscript must be printed at 1.5 spacing in a single column on one side of A4 paper. All pages should be numbered. Length of paper should be between 4,000 to 6,000 words (8 to 12 pages), excluding tables and illustrations. Metric units should be used and all non-standard symbols, abbreviations and acronyms must be defined.

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Abstract in both Malay and English, each in one paragraph and should not exceed 300 words. It should clearly identify the subject matter, results obtained, interpretations discussed and conclusions reached.

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Chapter of books and Symposium volumes:

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Article in Malay:

Lim, C.H. & Mohd. Shafeea Leman, 1994. The occurrence of Lambir Formation in Ulu Bok Syncline, North Sarawak. *Geol. Soc. Malaysia Bull.*, 35:1-5. (in Malay with English abstract)

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